1	The Nhecolândia wetland: Natural and anthropogenic influences on south-east
2	Pantanal, Brazil.
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Abstract: The Nhecolândia region covers the southern wetlands of the Pantanal 15 basin in Brazil. Characterised by myriad shallow freshwater and alkaline-saline lakes,, 16 the distinct natural features of the Nhecolândia wetland make it highly sensitive to 17 18 climate change and human impact. This study summarises the natural and social 19 aspects that have, and may potentially cause changes to this delicate wetland. Here, 20 we analyse the response of the wetland to historical changes in climate and human activity and use this understanding to forecast responses to future changes. The data 21 presented here show that this region is particularly sensitive to alterations in the flood 22 regime, droughts and deforestation which are intrinsically related to global and local 23 climate changes and the intensification of the cattle ranching activities, which include 24 25 deforestation and the introduction of cultivated pastures.

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29 The Pantanal basin, located in central-west Brazil, is one of the world's largest wetlands, covering more than 140,000 km². Due to the low relief of the entire 30 basin (ranging from 80 to 200 m in elevation) and a complex hydrological network of 31 rivers and lakes, large tracts of the basin are susceptible to seasonal floods that can 32 last from 4 to 8 months in some areas (Assine 2015; Marengo et al. 2016b). Acting as 33 a wetland mosaic, the Pantanal is characterised by different climatic, hydrologic and 34 topographic gradients (Fig. 1; McGlue et al. 2017). During the dry season, from April 35 to September, the waters remain in perennial rivers and lakes, while extensive and 36 continuous grasslands are exposed in between the "islands" of vegetation where 37 cerrado savanna-type forest predominates (see Biodiversity section). Due to its 38 39 geomorphology and hydrology, the Pantanal acts as a natural water reservoir, filled up during the rainy season and slowly being drained out during the dry season, 40 helping to keep the water balance of the south Paraguay and Paraná rivers basins 41 (Paz et al. 2014; Marengo et al. 2016b). 42

43 The Nhecolândia region encompasses the southern wetlands of the Pantanal basin (Fig. 1). This region is located to the south of the Taquari Megafan, 44 where myriad shallow lakes have developed in an inactive lobe of the fan, with at 45 least two distinct chemical compositions: fresh-water and alkaline-saline (McGlue et 46 al. 2017; Gerreiro et al. 2017; see The Nhecolândia region geomorphology section). 47 This leads to a unique environment, with cerrado forests, native pastures and 48 49 herbaceous vegetation around the lakes (Bergier et al. 2016). The high 50 evapotranspiration rate, which exceeds the annual rainfall (Marengo et al. 2016b), 51 contributes to an extended dry season that drastically reduces the number of filled 52 lakes in the area and increases the salinity of the alkaline-saline lakes (McGlue et al. 2017). The presence of freshwater in the lakes of Nhecolandia, and the wider flood 53 dynamics of the region, have a direct impact on the economy. Not only do they 54 determine the availability of freshwater for humans and animals, but in so doing they 55 define the biodiversity in the area and the land use, impacting the distribution of 56 suitable areas for cattle ranching (Batista et al. 2012). 57

58 The sensitivity of the region means that anthropogenic climate and land 59 use change could have a significant impact in the Nhecolândia region. Projections of 60 changes in the occurrence of extreme floods and droughts may alter the ecological dynamic of the region. Human activities, predominantly cattle ranching, may also
impact the natural environment, modifying the ecosystems directly and indirectly, with
Bergier *et al.* (2013) and Marengo *et al.* (2016) concluding that natural or anthropic
modification in the hydraulic conditions can threat the Pantanal ecological balance.

Here, we present a review of the knowledge about the sensitive Nhecolândia region. The objectives of this study are: (1) to review the natural and social characteristics of the region, (2) to identify its potential sensitivity pinch-points and (3) to assess future scenarios for the evolution of the Nhecolândia wetlands.

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70 The Nhecolândia environment

71 Climate and hydrology

72 Paleoclimate of Nhecolândia

73 The sedimentary record of the Taquari Megafan indicates that, since the 74 Late Pleistocene, the depositional system was controlled by alternating humid and dry periods (Claperton 1993; Assine & Soares 2004; Whitney et al. 2011; McGlue et 75 al. 2017). Distinct cold and dry periods, between 45 and 19.5 kyr BP are described by 76 Whitney et al. (2011), with arid periods, in comparison with the present day climate, 77 78 remaining until 12.8 kyr BP and occurring again in between 10 and 3 kyr BP. The 79 strongest modification of the landscape, with exposure and reworking of the sediments in arid conditions, occurred during these dry periods, coupled to the 80 associated changes on the rainfall amount (McGlue et al. 2016; 2017). 81

82 During the Early Holocene, the climate was predominantly dry when compared to present days (Whitney et al. 2011; McGlue et al. 2012; McGlue et al. 83 2017), including an intense drought episode (approx. 5,300-2,600 yr BP) interpreted 84 from a hiatus in sedimentation that affected the region (McGlue et al. 2012). A later 85 transgression coincided with increased flood and a more humid climate after that. 86 The development of the Nhecolândia lakes (see The Nhecolândia region 87 geomorphology section) is related to an abandoned lobe of the Taguari Megafan, with 88 89 sediments reworked and morphology reshaped by enhanced wind conditions during 90 more arid periods, leading to an aeolic and evaporative environment, with increasing salinity in the hydrological isolated lakes since ~910 yr BP (Guerreiro *et al.* 2017;
McGlue *et al.* 2017).

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Present climate and hydrology of Nhecolândia

95 The geographical position of the Pantanal basin, in the centre of the South American continent and south of the Amazonas basin, means its climate is 96 modulated by two phenomena: the Intertropical Convergence Zone (ITCZ) and the 97 South Atlantic Convergence Zone (SACZ; Lenters & Cook 1999; Marengo et al. 2001; 98 Marengo et al. 2016a). Being modulated by both the ITCZ and the SACZ, the South 99 American Monsoon System delivers concentrated rains from October to March (100 100 to 300 mm/month) and a dry period during the rest of the year (0 to 100 mm/month; 101 Garcia 1984; Vera et al. 2006; Marengo et al. 2016a). 102

103 Despite the Pantanal's position within the tropical latitudes, its climate is 104 closer to semi-arid than to tropical conditions (Marengo et al. 2016b). With an average air temperature of 24°C and an annual rainfall of 1,100 mm, the climate of 105 Pantanal presents high evapotranspiration rates of 1,400 mm (Furian et al. 2013; 106 Marengo et al. 2016b; McGlue et al. 2017), and an interannual variability that can 107 cause intense dryness or drastic floods (Marengo et al. 2016a). The annual 108 109 monsoonal rainfall patterns are affected by large-scale climate phenomena such as 110 the El Niño Southern Oscillation, and the Atlantic Ocean variability. Furthermore, regional scale phenomena, such as regional water balance, soil wetness and soil 111 moisture storage influence the climate of the Pantanal (Bergier 2010; Marengo et al. 112 2016b; Bergier et al. 2018). 113

Throughout the year, the Pantanal undergoes a wide range of temperature variation, with daily contrasts between 1°C and 40°C common in winter (June-August). This is partly due to the geomorphological interference of the Andes (west) and the Brazilian Central Plateau (east), with atmospheric weather patterns creating an air corridor capable of driving the southern polar air masses form northern Argentina directly to the Pantanal (Marengo et al. 2016a), in a phenomenon called *friagem*. 121 The drainage system of the Pantanal basin has approximately 500,000 122 km² of channels, with the main arteries being the Paraguay, Cuiabá, Piguiri, Taguari, Miranda and Negro rivers. The relief of the basin is minimal, which makes it 123 susceptible to a seasonal flooding pulse (Hamilton et al. 2002; Paz et al. 2014), 124 influenced by the Amazonian rainfall regime in the northern Pantanal rivers, which 125 126 propagates from north to south, only reaching the southern regions of Pantanal during the dry season (Por 1995; Hamilton et al. 2002; Fantin-Cruz et al. 2011; Paz et 127 al. 2014; Bergier et al. 2018). 128

The patterns of the rainfall in the Nhecolândia region display a 129 130 non-uniform distribution, with an east-west gradient evident (92 mm/month in the east compared to 74 mm/month in the west). In the Paraguay river, west of Nhecolândia, 131 132 the rainfall and the flood peak are disconnected, with the seasonal rainfall occurring ~ 4 months before the floods (Fig. 2; Por 1995; Hamilton et al. 2002; Fantin-Cruz et al. 133 2011) with notable influence of the Amazon rainforest in the moisture distribution 134 (Bergier et al. 2018). There is, therefore, a very complex relationship at different 135 136 scales between the hydrology of this region and the climatic patterns which drive it.

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138 Geological context

The Pantanal basin is positioned in west-central Brazil (with some portions of the region extending into Bolivia and Paraguay) with a surface area of approximately 140,000 km² (Fig. 1; Almeida & Lima 1959; Assine 2015; Assine *et al.* 2015). The basin is elongated (north-south), with borders controlled by normal faults and a maximum sedimentary thickness of 500 m (Weyler 1962; Assine 2015; Assine *et al.* 2015). The basement consists of low-grade Neoproterozoic metamorphic rocks of the Paraguay Belt (Alvarenga *et al.* 2009; Assine *et al.* 2015).

The origin of the Pantanal basin is related to the tectonic reactivation of the forebulge of the Andean orogeny, during the Gelasian stage (around 2.5 Ma ago, during the Early Pleistocene, Quaternary), as part of the Andean foreland system (Horton & DeCelles 1997; Ussami *et al.* 1999; Assine *et al.* 2016). However, the origins of the structural settings of the basin are related to a regional epeirogenic uplift occurring during the Early Pleistocene, when the basin major faults were generated and when the basin depocenter started to subside (Almeida *et al* 2000;
Assine 2015; Assine *et al.* 2015; Assine *et al.* 2016).

The infilling of the basin is characterised mainly by siliciclastic 154 sediments (Weyler 1962; Assine et al. 2016). The main sequences show 155 156 fining-upward trends, from conglomerates and coarse-grained sandstones at the base, to fine and medium sands, locally coarse-grained, in the upper layers. The 157 mineral composition is mainly quartz, with some fragments of metamorphic basement 158 159 rocks related with the coarse portions (Weyler 1962; Assine 2015; Assine et al. 2015). 160 Due to the process of subsidence, related to normal faults, alluvial sedimentary 161 systems occur in the east border of the basin, grading to alluvial fans that occupy the major part of the basin, and to fluvio-lacustrine systems related to the Paraguay River 162 (Assine et al. 2015). The alluvial fans are the most representative sedimentary 163 systems, comprising the São Lourenço, Aquidauana and Taquari megafans (Zani et al. 164 165 2012; Assine et al. 2015).

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167 Geomorphology

168 The Taquari Megafan

The Nhecolândia region forms the southernmost part of the Taquari 169 170 Megafan (Fig. 1), which has been depositing sediments in the Pantanal tectonic 171 depression since the Early Pleistocene (Horton & DeCelles 1997; Ussami et al. 1999; Assine & Soares 2004; Assine 2015; Assine et al. 2015). The 49,000 km² alluvial fan 172 deposits extending for ~250 km from apex (190 m a.s.l.) to toe (85 m a.s.l.), with a 173 circular geometry suggests a multi-lobed evolutionary history (Assine & Soares 2004; 174 175 Assine 2015; Assine et al. 2015; Zani et al. 2012). Characterised by low gradient 176 slopes, a large number of meandering, single channel rivers extend across the fan surface. The high levels of fluvial activity across the fan, including frequent channel 177 avulsions (Assine 2005; Assine 2015; Assine et al. 2015; Buehler et al. 2011), mean 178 that relic paleo-channels are common-place. During the flood season these channels 179 are activated and convey water and sediment across the fan, resulting in a 180 highly-connected landscape which is less reliant on active channel systems to 181 182 distribute sediment, nutrients and water.

183 The Taquari Megafan may be delineated into two distinct 184 geomorphological zones (Souza et al. 2002; Assine & Soares 2004; Buehler et al. 185 2011; Assine et al. 2015); the upper fan displays a confined meander belt incised into abandoned depositional lobes, whilst in the lower fan the river is straighter and is an 186 active depositional environment (Fig. 3). The currently active depositional region is 187 188 highly dynamic with both partial crevasse-splays and complete avulsions (Slingerland & Smith 2008; Buehler et al. 2011) occurring across the lobe. Assine (2005) and 189 Buehler et al. (2011) describes a complete avulsion that occurred between 1988 -190 191 1999 at the fan toe, with the resulting distributary channel conveying 70% of the discharge between 1995 - 1997 (Padovani et al. 2001). Recent sedimentation in the 192 193 upper reaches of the depositional fan lobe has resulted in increasing overbank flow 194 and the development of multiple crevasse-splay deposits along the main channel on 195 the northern most side of the fan. Assine (2005) speculates that this rapid sedimentation may result in a major avulsion, potentially initiating a new 196 depositionary lobe. 197

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The Nhecolândia region geomorphology

The Nhecolândia region covers the southernmost 25,000 km² of the 200 201 Taquari Megafan (Marengo et al. 2016b). Unique to the Nhecolandia region is the 202 coexistence of thousands of fresh-water ('baías') and alkaline-saline ('salinas') lakes 203 (Fig. 4; Furguim et al. 2008; Evans & Costa 2013; McGlue et al. 2017; Guerreiro et al. 2017). The baías may be up to two meters deep and are found in various areal 204 extents. During the flood season baías may coalesce and form temporary flood 205 206 channels, or 'vazantes', which recede back to isolated ponds during the dry season. 207 There is, therefore, much connectivity between baías. The salinas, in contrast, are 208 generally 500 - 1,000 m in diameter, 2 - 3 m deep (Barbiero et al. 2008; Furian et al. 209 2013; McGlue et al. 2017; Guerreiro et al. 2017) and often raised above the 210 surrounding ground level. They are often enclosed by sand embankments, or 'cordilheiras' which are covered by dense savannah vegetation, and so remain 211 isolated from the flood waters. This maintains distinct geochemical signatures 212 between different salinas (see Soils and hydrochemistry of lakes section). Costa and 213 Telmer (2006; 2007) also distinguish an intermediate type of lake, 'doces', which they 214

215 define as lakes in transition from the fresh-water '*baias*' to the more brackish 216 '*salinas*'.

The origin of these different lake types is not overtly clear. Early 217 hypotheses suggested an aeolian explanation with the lakes forming in troughs and 218 219 hollows produced by aeolian deflation (Tricart 1982). Klammer (1982) identified a series of longitudinal fossil dunes aligned in the NNE and NNW directions and 220 concluded that the salinas formed as salt pans due to ponding in the interdune areas. 221 Assine & Soares (2004) associate the sand ridges, or 'cordilherias', with relict 222 223 aeolian lunette sand dunes, although they found no evidence of the longitudinal 224 dunes interpreted by Klammer (1982). Barbiero et al. (2002) suggested that the 225 salinas are, in fact, more recent products of the Pantanal ecosystem, forming over decades rather than centuries of millennia, highlighting their importance in the 226 modern ecosystems of the Pantanal and their response to human and climatic 227 pressures (Costa & Telmer 2006). The relatively inactive nature of this part of the 228 Taquari Megafan means that large scale geomorphic reworking of sediments as a 229 230 result of channel avulsion provides the lakes of the Nhecolandia region with the large timescales needed to form and maintain distinct geochemical signatures in the 231 isolated salinas (Becker et al. 2017; Guerreiro et al. 2017; McGlue et al 2017). This 232 may help to explain why these landforms are not found on the more active northern 233 234 and western areas of the Taquari Megafan.

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Soils and hydrochemistry of lakes

237 Much research has dealt with the chemical variability in the lakes of the 238 Nhecolândia region and the related soils forming within and around those lakes 239 (Barbiero et al. 2002, 2008; Soares et al. 2003; Furquim et al. 2008, 2010; Almeida et 240 al. 2009; Barros Nogueira et al. 2011; Rezende Filho et al. 2012; Furian et al. 2013; Bergier et al. 2016; McGlue et al. 2012, 2017, among others). The co-existence of 241 alkaline-saline and freshwater lakes is attributed by Barbiero et al. (2008) to a 242 243 differential hydrological regime whereby water is lost from the saline lakes by evaporation, and an inflow coming from the freshwater system is directed towards the 244 depressions of the non-saline lakes, refilling them (Fig. 4). It is commonly agreed that 245 246 a large part of the chemical variability of the Nhecolândia lakes is a direct product of

the evaporation processes (Furquim 2007; Furquim *et al.* 2008; Rezende Filho *et al.*2012, McGlue *et al.*, 2012, 2017), that in turn produces the precipitation of authigenic
mineral phases (Furquim *et al.* 2010).

However, Rezende Filho *et al.* (2012) question evaporation as the main factor in the salinisation of the ponds, given the current climatic conditions in the area. In fact, McGlue et al. (2012) describe a drought stage in the Holocene (5,300-2,600 yr BP) followed by a sudden increase in the flooding pulses, related to a more humid part of the Holocene. Therefore, it seems somehow contradictory to propose evaporation as the main reason for the presence of *salinas* in Pantanal (McGlue et al., 2017) as evaporation rates were likely lower during this formative period.

257 As a potential explanation, Furian et al. (2013) suggest that under poor drainage conditions, salinity can be maintained naturally in the landscape despite 258 259 relatively humid mean climatic conditions, meaning that a past arid climate would not be necessary to produce the presently saline lakes in Nhecolândia area. They 260 propose the ponds are filled by groundwater during the dry season, and the outflow 261 262 would be prevented by the chemically cemented soils (Fig. 4). The results obtained by Guerreiro et al. (2017) fully support the idea of alkaline-saline lakes not being a 263 relict of past aridity events. McGlue et al. (2012, 2017) also support Furian et al. 264 (2013) explanation, adding the presence of cordilheiras as an important 265 266 geomorphological feature acting as barriers to the floods and contributing to the 267 evaporation and high salinity of the water in the salinas.

268 Biodiversity

The biodiversity of Pantanal wetland secures its importance as one of the largest freshwater ecosystems in the world. Affected by the highly variable water regime determined by the contrasts between the wet and the dry seasons, two very different styles of vegetation coexist (Alho 2008): mesic (related to a moderate amount of humidity) and xeric (needing a small amount of humidity to survive).

The flora of Pantanal has been studied in detail by several authors (Pott & Pott 1996, 2009; Pott *et al.* 2011a, 2011b; Scremin-Dias *et al.* 2011), and is characterised by its mosaic-like heterogeneity (Adámoli 1982), including inundated floating plants, seasonal flooding fields, gallery forests, scrub and semi-deciduous forests and different kinds of Cerrado savanna (Alho 2005, 2011). In the Nhecolândia
region there are more detailed studies on the forest vegetation (Ratter *et al.* 1988).
These authors imply that the vegetation in Nhecolândia presents similarities to the
deciduous and semideciduous forests of Central Brazil.

282 There are two main biomes in the Pantanal wetland (Alho et al. 2000; Alho 2005; Sabino & Prado 2006): the Cerrado and the Pantanal biomes. The 283 Cerrado biome is a savanna-type forest, including open savanna (grasslands) with 284 285 sparse trees and closed woodlands with little grass. The Pantanal biome, restricted to 286 the west of the region, corresponds to the floodplain areas with mainly wetland-type 287 (Chaco and Amazonian) vegetation, including aguatic macrophytes among other plants. Although in the Brazilian Pantanal there are various degrees of Cerrado cover, 288 in the region of Nhecolândia, it consists of between 40 and 50% of the total 289 290 vegetation.

Regarding the wildlife, the feeding and breeding conditions are directly 291 292 linked to the hydrological cycle and may be divided in four stages (Alho 2008, 2011): 293 flooding, flood season, drainage season and dry season. The water availability affects not only the vegetation, but also the niches and the nutrient cycles for the 294 295 animals that live in Pantanal. Moreover, many endangered species can be found in the Pantanal such as the jaguar and the marsh deer (Alho et al. 2003). Approximately 296 297 113 reptile species live in the Pantanal and 189 in the Cerrado biome. Specifically, 298 amphibians show low diversity but high abundance in the area, with some 80 species occurring on the upper Paraguay Basin (Alho, 2008). 463 bird species have been 299 recorded only for the floodplains, increasing to 665 species when the uplands are 300 included, and 837 species when including the Cerrado biome (Alho et al. 2000, 2008; 301 302 Alho 2005; Sabino & Prado 2006; Tubelis & Tomás 2002; Silva 1995). Mammals are also quite diverse, with up to 195 species identified so far (Alho et al. 2003, 2011; 303 304 Alho & Gonçalves 2005; ANA 2004). Finally, Alho (2008) provides an estimation of 400 fish species living in the Pantanal region (including plateaus and plains). 305 However, although there are many studies on the biodiversity of Pantanal as a whole, 306 only a few show the close links between the annual floods and the wildlife and focus 307 308 specifically on the Nhecolândia region. (e.g. Mamede & Alho 2006). Taking into account that floods directly influence the habitats of the Pantanal fauna and flora, it is 309

310 important to fully understand the wetland evolution in order to predict changes that

- 311 may affect its rich biodiversity in the future.
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313 Anthropogenic influence on Nhecolândia

314 Prehistorical occupation

The human occupation in the Pantanal started around 27,000 years BP 315 with a settlement located 100 km north of the basin (Oliveira 1994; Schimitz et al. 316 317 2009; Vialou et al. 2017). Archaeological evidence suggests the occupation of the rest of the basin began ~8,000 years ago in locations close to the Paraguay river 318 (Schimitz et al. 2009) and, later in the western area of Nhecolândia, in a coincident 319 period of a shift towards wetter conditions in the basin (McGlue et al. 2012). Some 320 ceramic sites related to natural carbonate mounds (4,400 years B. P.; Oliveira 1994; 321 322 Schimitz et al. 2001; Oliveira et al. 2017) have been identified in a region called 323 Pantanal do Miranda (Fig. 1).

324 Although there are no known archaeological sites in the centre of the 325 Nhecolândia region, the west and south-west borders are full of archaeological sites, specially related to the Pantanal do Miranda mounds, geomorphological features 326 327 resulting from the differential erosion of carbonate paleolakes and their surroundings 328 (Oliveira et al. 2017). These mounds represent elevation highpoints on the Pantanal 329 plain and are constantly exposed during flooding, forming "vegetation islands" that are rich in ceramic fragments of the Pantanal Tradition (Oliveira 1994; Schimitz et al. 330 2001, 2009). The people that occupied these mounds had a close relation with the 331 surrounding floodplain environment, with small settlements and a fish based diet 332 333 (Schimitz et al. 2001).

The carbonate mounds of Pantanal do Miranda are a more stable and elevated terrain, with denser vegetation, and closer to the river than the sand elevations (*cordilheiras*) of Nhecolândia, making them a more suitable place to live for ancient humans (Oliveira 1994; Schimitz *et al.* 2001, 2009). This is probably the main reason for the differential preservation of human remains in these settings.

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Modern occupation

341 Despite the presence of some native ancient human groups in Pantanal, some native modern groups from the Amazon migrated to the region prior to the 342 arrival of the Europeans (Oliveira 1996) in the fifteenth century. The discovery of this 343 "New World" marked a step-change in the occupation of the Pantanal. European 344 occupation was fast, with the north part of Pantanal being first occupied by the 345 346 Europeans in the 1520s, as the rush for silver mining, earlier found in the nearby Andean region, spread to the Pantanal (Oliveira 1996). However, as the geological 347 characteristics of Pantanal and its surroundings do not favour the development of 348 349 abundant precious metals, with only some gold occurrences being discovered in the 350 1800's (Oliveira 1996), this occupational period was short lived.

351 The region became permanently occupied in the early 1900s, with an extensive cattle 352 ranching industry expanding to fill the demand for dry beef. Population in the region grew more consistently after the Brazilian government colonization programs 353 between 1960 and 1980 (Coutininho 2005; Abreu 2010). During this period changes 354 in land use were extensive, and cattle ranching became the main economic activity in 355 356 Nhecolândia. The current estimated population of Nhecolândia region is ~3,000 357 people (Abreu 2010). Therefore, the habitation of this region is relatively new and conditioned to the hydrological and climatological characteristics of the past century. 358 Inbuilt societal resilience to long term variations in hydrology and climate may 359 therefore be lacking, meaning that the communities in this region are likely highly 360 sensitive to potential future changes. 361

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Cattle ranching in Nhecolândia

The main characteristic of the cattle ranching industry in the Nhecolândia region is the cattle graze system (transhumance), with the farmer driving the livestock through different farm lands during the year, to avoid the seasonal floods and to reduce the stress of the grazing lands in the dry season (Rosa *et al.* 2007). The production is focused on calf rearing, independent of the environmental conditions, and the total number of animals is around 1 million (Batista *et al.* 2012).

The cattle production in this region occurs in extensive farms due to the need of native pastures, with approximately 8 pastures per farm (Rosa *et al.* 2007). 372 The land should include variations in relief, covering low lands, such as baias and 373 vazantes, and high lands, such as cordilheiras. The pastures are essentially native, 374 but due to production limits, some cultivated pastures are being introduced, especially in the flood free areas (Pott 1982). This involves the removal of native 375 vegetation and the introduction of some invasive and non-native species, having a 376 377 dramatic impact on the local ecosystems. Sometimes it is necessary to synchronize 378 the movement of the cattle with other farms, in case of intense floods or drought, in an integrated management system that involves the farmers (Rosa et al. 2007; Abreu 379 380 et al. 2010). Transhumance, a practice inherited from the past century, may not be 381 resilient to future changes. It is a crucial component of Nhecolandia economical 382 activity, and it depends mainly on the availability of pastures for the cattle to feed on. Thus, hydrological changes in the flood regime and in the extension of land affected 383 by the floods are expected to have a direct impact on this activity, potentially reducing 384 the primary food source for the animals and therefore forcing changes in the current 385 transhumance patterns. 386

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The impacts of human activity

There is no record of prehistorical alterations in the landscape, due to 389 390 the hunter-gather character of all native groups that lived in or close to Nhecolândia (Schimitz et al. 2001; Bespalez 2015). The absence of economically valuable natural 391 392 resources, such as minerals, prevented the area from human impact until the beginning of the cattle ranching activity (Oliveira 1996). The main modification is the 393 removal of native vegetation and/or the native pastures in substitution for cultivated 394 395 ones (Harris et al. 2005). In between 2003 and 2010, 12% of the forested area were replaced by cultivated pastures (Paranhos Filho et al. 2014). The deforestation of the 396 397 region leads to desertification due to the changes in the local water cycle (i.e. 398 humidity loss), especially around the saline lakes, where specialised vegetation has evolved (Sakamoto et al. 2012). The constant growth of the cattle ranching activity in 399 recent years is also leading to an increase in stress of the native vegetation due to 400 401 the use of pesticides and herbicides and the burning of weeds in some areas (Harris et al. 2005). The influence of some anthropic modifications, such as unsustainable 402 agriculture in river head waters, can be observed in distinct areas, inducing variations 403

in water volume and affecting the sediments and nutrients carried by the rivers from
the high lands of Pantanal to the low lands, such as Nhecolândia (Bergier *et al.* 2013).
Although these impacts may be recent, they are likely to be major drivers of
environmental and ecosystem change in the near future of this important wetland
region.

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410 Future perspectives

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Nhecolândia's sensitivity to change

The sensitivity of the Nhecolândia region to environmental modifications can be considered high since it is now often classified as a semi-arid region (Marengo et al. 2016b), were the annual rainfall is lower than the annual evapotranspiration rate. These factors, combined with the concentration of rains (October to April), can leave the region without rainfall for 5 to 7 months (Marengo et al. 2016b).

The waters that characterise Nhecolândia as a wetland are related to 418 the water regimes of Negro and Paraguay rivers, both with distinct flooding cycles. 419 420 The Negro river is directly related to regional rainfall and the water-table regime, acting as a tributary of Paraguay river. The Paraguay river has its origin in the 421 422 Amazon basin and it is highly affected by the Amazon region climate (Por 1995; 423 Hamilton et al. 2002; Fantin-Cruz et al. 2011; Bergier et al. 2018), which by extension 424 affects the Nhecolandia region. Therefore, the hydrological regime of the region may be greatly affected by modifications to the hydrological regime of the Negro and 425 426 Paraguay rivers as well as by changes to the precipitation regime of the Nhecolândia 427 region itself. The sensitivity in the variation of the water level of the Paraguay river, 428 observed by Collischonn et al. (2001), shows that some droughts can last as long as a decade, as observed between 1962 and 1973. Conversely, the subsequent 429 period(from the mid-1970's to 2001) shows higher water levels. This can be related 430 to the increased river discharge due to deforestation for new pasture areas (Silva et 431 al. 2011). 432

The connection between the Amazonian and Pantanal environments is highlighted by Bergier *et al.* (2018) who show that the deforestation in the Amazon rainforest directly affects the distribution and intensity of the rainfall in Pantanal; which can cause natural disasters such as severe floods and droughts. Human alterations, as well as changes in climate, in these basins are likely to be compounded in the Nhecolândia region, where complex interactions between multiple source zones create a dynamic, ecological hotspot.

With respect to the ecosystems supported by the Nhecolândia region, 440 Alho & Silva (2012) report the sensitivity of some animal groups in the Nhecolândia 441 442 ecosystem to climatic variations, where variations in flooding frequency and 443 prolonged periods of drought have been shown to impact upon community structure, 444 breeding habits, population size and behavioural ecology of animals. This is due to the adaptation of these communities to the seasonal regime of the region (Mamede & 445 Alho 2006). This adaptation means that future changes in the hydrological regime of 446 the region may have devastating impacts on the ecology and biodiversity of the 447 region. Potential future changes may be intensified by deforestation and/or 448 substitution of native vegetation and pastures for cultivated pastures (see Cattle 449 450 ranching in Nhecolândia and The impacts of human activity sections). Additionally, these changes will reduce habitats upon which many local food-webs rely (Sakamoto 451 et al. 2012). The process of deforestation can be a direct anthropogenic action or a 452 consequence of expanding cattle ranching activity, which induces modifications in the 453 454 farms landscape, soil and vegetation that increases with the number of animals 455 (Sakamoto et al. 2012; Rodela et al. 2007). From a social aspect, the modernisation 456 of cattle ranching techniques and environments can negatively affect the cultural stability of the Pantanal population, changing the way they live in relation to the 457 458 original environment (Rosseto 2003).

459 460

Lessons from other wetland environments

Given the sensitivity of the Pantanal wetland to changes in climate and the relatively limited anthropogenic influence in the region, it may be pertinent to assess records of change from other ancient and modern wetland environments around the World (e.g. Pla-Pueyo et al. 2015) in order to glean information as to how the Pantanal wetland may respond in the future. Looking at modern wetlands in the lbera wetland, north Argentina, analysis of macro-fauna suggested that extensive dry 467 periods provide refuges to animals with high dispersal ability, despite being a problem 468 to animals with low dispersal ability (Ubeda *et al.* 2013). It has also been shown that 469 the mortality of large mammals increased during dry periods due to pollution and 470 cyanobacteria dominance in small water bodies (Ubeda *et al.* 2013).

471 More locally, Barros & Albernaz (2014) show that modifications in 472 climate in an Amazonian wetland, can cause intense droughts or inundations, which in both cases can lead to environmental change, causing losses within the specific 473 474 environment of the wetland. They highlight that sensitivity to climate changes is intrinsic to wetlands, due to their relation with both dry and wet environments. The 475 476 results of these modifications will challenge the adaptability of species (composition 477 and distribution) with consequences for the related human population (Barros & Albernaz 2014). 478

479 480

Sustainability in Nhecolândia

Projections of future climate change in the region show that temperature may increase up to 7°C in the next hundred years (Eta-HadGEM2 ES model; Marengo et al. 2016b), leading to an increase in evaporation and a decrease in rainfall values (Marengo et al. 2016a). The evolution of the current climate into a more arid one may affect all aspects in the Nhecolândia region, from the rainfall and flood frequency to the sedimentation rates of the Taquari Megafan, leading to drastic changes in the landscape, water availability and ecosystems.

488 There are many studies in the literature discussing how aridification has been recorded in the fossil record and how it affects the morphology and 489 490 sedimentation style of alluvial fans that were previously reflecting wetter conditions (Bull & Schick 1979; Wells et al. 1987; Harvey & Wells 1994; Pla-Pueyo et al. 2015). 491 Blair and McPherson (1994) related alluvial fan aggradation in the Mojave Desert of 492 the southern USA to transitions from a wetter to a dryer climate, helped by ephemeral 493 processes such as storms. Sheets et al. (2002) demonstrated through an experiment 494 that most of the deposition in an alluvial fan occurs as a result of short-lived flows 495 caused mainly by overbank spills, flow expansions and failed avulsions. Highlighting 496 497 the importance of discrete single events, rather than prolonged wet/dry cycles on alluvial fan evolution. Therefore, potential changes towards more infrequent, butmore intense rainfall may be beneficial for the growth of the Taquari Megafan.

500 If the Taquari Megafan were to undergo an aridification process, a 501 decrease in the effective soil moisture would reduce the vegetative cover on the 502 hillslopes, which would in turn result in an increase in soil erosion and a concomitant increase in sediment supply donwnstream (Harvey & Wells 1994; Harvey et al. 1999; 503 Silva et al. 2011) in the floodplain area. This area is currently occupied by cattle 504 505 ranching pastures. Moreover, in such a sensitive system, the decrease of moisture in 506 the soil, together with the change in temperatures, would result in the reduction and 507 potential disappearance of the deciduous forests and most of the aquatic macrophytes in the area, highly impacting the wildlife depending on such vegetation. 508 The landscape would change, as would the current ecological niches with it. Under 509 these future environmental conditions, the flood regime would be expected to change 510 511 according to the reduced rainfall and the changes in the watershed soil cover, leading to a shortening of the wet season, and more frequent flash-flood events substituting 512 513 channelised, less flashy floods. With an increased evapotranspiration, a decreased 514 rainfall and floods more concentrated in certain periods of the year, periods of drought are more likely to occur. This would result in some of the lakes drying out 515 permanently, and the remaining ones concentrating; with, the freshwater ones 516 becoming more saline, drastically reducing the amount of freshwater availability in 517 518 the remaining lakes.

519 Given the above, climate change has the potential to be extremely 520 damaging to the current biodiversity of the Nhecolândia region, and would be expected to have a destructive impact both on the wild ecosystems and human 521 522 activities. In the meantime, the Pantanal is currently suffering anthropogenic 523 modifications of its natural landscape, such as deforestation and/or substitution of the 524 natural pastures by introduced cultivated species that leads to the competition of the 525 native and exotic species, degradation of habitats and exposure of the sandy soil (Alho & Sabino 2011). The sandy soil is not suited for the growing of pastures and 526 therefore, reduces the feeding areas for cattle. These modifications can accelerate 527 the process of regional climate change, turning Nhecolandia into an arid place, with 528 dramatic consequences for flora, fauna and human occupation. 529

530

531 Conclusions

532 The Pantanal region is a mosaic representing millennia of climatic 533 changes, displaying a landscape that is the result of multiple shifts from humid to more arid environments. As such, the current landscape geomorphology, ecology and 534 productivity of this composite landscape reflects its unique position at the boundary of 535 several key atmospheric circulatory patterns (e.g. the ITCZ). Within the Pantanal, the 536 Nhecolândia region remains one of the most understudied, yet sensitive area of the 537 538 basin. The unique hydrological regime of this region, currently classified as semi-arid, means the sensitivity of the Nhecolândia sub-region to environmental modifications 539 can be considered high. The water regimes of the main rivers are distinct; related to 540 regional rainfall (Negro river) and Amazon region climate cycles (Paraguay river). In 541 542 addition, human modifications and pressures also increase the sensitivity of the region to climate change. The main human impact in the region is deforestation in 543 544 order to introduce cultivated pastures .The resultant changes in habitats, land cover 545 and hydrological regimes mean that the entire Nhecolândia system is impacted. . Some native animal groups in the Nhecolândia ecosystem are highly sensitive to 546 climatic variations, so any small climatic shift may produce an impact upon 547 community structure, breeding habits, population size and behavioural ecology. 548 549 Therefore, any future changes to the climate, land cover or ecosystem of the 550 Nhecolândia region will result in complex reactions which will impact the society and ecology of the entire Pantanal basin. 551

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553 References

554

555 Abreu U. G. P., McManus C. & Santos S. A. 2010. Cattle ranching, conservation and 556 transhumance in the Brazilian Pantanal. *Pastoralism*, **1**(1), 99-114.

557

Adámoli J. 1982. O Pantanal e suas relacoes fitogeográficas com os Cerrados:
discussao sobre o conceito de "Complexo do Pantanal". *In: Anais XXXII Congresso Nacional de Botânica, Teresina, Brazil, 1981,* 109-119.

Alho C. J. R., Strüssmann C., & Vasconcellos L. A. S. 2000. Indicadores da 562 magnitude da diversidade e abundância de vertebrados silvestres do Pantanal num 563 mosaico de hábitats sazonais. In: Anais do III Simpósio sobre Recursos Naturais e 564 565 Sócio-Econômicos do Pantanal, 2000, 1-54. 566 Alho C. J. R., Strussmann C., Volpe M., Sonoda F., Marques A. A. B., Schneider M., 567 Santos-Junior T. S. & Margues S. R. 2003. Conservação da Biodiversidade da Bacia 568 do Alto Paraguai. UNIDERP, Campo Grande. 569 570 Alho C. J. R. 2005. Ecology and Conservation. In: Fraser L.H. & Keddy P.A. (eds.) 571 The Pantanal. In The World's Largest Wetlands. Cambridge University Press, New 572 573 York, pp. 203-271. 574 575 Alho C. J. R. & Gonçalves H. C. 2005. Biodiversidade Do Pantanal. Ecologia E Conservação. UNIDERP, Campo Grande. 576 577 Alho C. J. R. 2008. Biodiversity of the Pantanal: response to seasonal flooding 578

regime and to environmental degradation. *Brazilian Journal of Biology*, **68**(4), 957-966.

581

561

Alho C. J. R. & Sabino, J. 2011. A conservation agenda for the Pantanal's biodiversity. *Brazilian Journal of Biology*, **71**(1), 327-335.

584

Alho C. J. & Silva, J. S. 2012. Effects of severe floods and droughts on wildlife of the Pantanal wetland (Brazil)—a review. *Animals*, **2**(4), 591-610.

587

588 Almeida F. F. M. 1945. *Geologia do sudoeste matogrossense*. Departamento
589 Nacional da Produção Mineral, Rio de Janeiro.

590

591 592	Almeida F. F. M. & Lima M. A. 1959. <i>Planalto centro-ocidental e pantanal Mato-Grossense</i> . Conselho Nacional de Geografia, Rio de Janeiro.	nal	
593			
594	Almeida F. F. M., Brito Neves B. B., & Carneiro C. D. R. 2000. The origin and	ind	
595 596	evolution of the South American Platform. <i>Earth-Science Reviews</i> , 50 (1), 77-111.		
590 597 598	Almeida T. I. R., Fernandes E., Mendes D., Branco F. C. & Sígolo J. B. 2007. Distribuição espacial de diferentes classes de Lagoas no Pantanal da Nhecolândia,		
599 600	MS, a partir de dados vetoriais e SRTM: uma contribuição ao estudo de sua compartimentação e gênese. <i>Geologia USP. Série Científica</i> , 7 (2), 95-107.		
601			
602603604605	Almeida T. I. R. D., Paranhos Filho A. C., Rocha M. M. D., Souza G. F. D., Sígolo J. B. & Bertolo R. A. 2009. Estudo sobre as diferenças de altimetria do nível da água de lagoas salinas e hipossalinas no Pantanal da Nhecolândia: um indicativo de funcionamento do mega sistema lacustre. <i>Geociências (São Paulo)</i> 28 , 401–415.		
606			
607 608 609	Almeida T. I. R., Calijuri M. D. C., <i>et al.</i> 2011. Biogeochemical processes and the diversity of Nhecolândia lakes, Brazil. <i>Anais da Academia Brasileira de Ciências</i> , 83 (2), 391-407.		
610			
611 612 613	Alvarenga C. J., Boggiani P. C., Babinski M., Dardenne M. A., Figueiredo M., Santos R. V., & Dantas E. L. 2009. The amazonian palaeocontinent. <i>Developments in Precambrian Geology</i> , 16 , 15-28.		
614			
615	ANA (AGÊNCIA NACIONAL DE ÁGUAS). 2004. Strategic Action Program for the		

616 Integrated Management of the Pantanal and the Upper Paraguay River Basin.

617 618 ANA/GEF/PNUMA/OEA, Brasília.

619 620	ANA (AGÊNCIA NACIONAL DE ÁGUAS). 2017. Base de dados da Agência Nacional de Águas ANA Brasília http://www.ana.gov.br	Base de dados da Agência Nacional	
621	de Aguas. Awa, Brashia, http://www.ana.gov.br	_ [
622 623 624	Andrades Filho C. O., Zani H. & Ribeiro B. M. G. 2016. Cronologia relativa de eventos deposicionais no megaleque do rio Taquari revelada por hipsometria. <i>Boletim Geográfico do Rio Grande do Sul</i> , 27 , 51-67.		
625 626 627	Assine M. L. & Soares, P. C. 2004. Quaternary of the Pantanal, west-central Brazil. <i>Quaternary International</i> , 114 , 23-34.		
628629630631	Assine M. L. 2005. River avulsions on the Taquari megafan, Pantanal wetland, Brazil. <i>Geomorphology</i> , 70 (3), 357-371.		
632 633 634	Assine M.L. 2015. Brazilian Pantanal: A Large Pristine Tropical Wetland. <i>In:</i> Vieira B., Salgado A., Santos L. (eds) <i>Landscapes and Landforms of Brazil.</i> World Geomorphological Landscapes, 5 , Springer, Dordrecht, 135-146.		
 635 636 637 638 639 	Assine M. L., Merino E. R., Nascimento Pupim F., Azevedo Macedo H. & Santos, M. G. M. 2015. Trato de sistemas aluviais do Quaternário da Bacia do Pantanal. <i>Brazilian Journal of Geology</i> , 45 (3), 475-489.		
 640 641 642 643 644 	Assine M. L., Merino E. R., Pupim F. N., Warren L. V., Guerreiro R. L. & McGlue M. M. 2016. Geology and Geomorphology of the Pantanal Basin. <i>In</i> : Bergier I. & Assine M. (eds) <i>Dynamics of the Pantanal Wetland in South America.</i> The Handbook of Environmental Chemistry, 37 , Springer, Cham, 23-50.		
645 646	Barros D. F. & Albernaz, A. L. M. 2014. Possible impacts of climate change on wetlands and its biota in the Brazilian Amazon. <i>Brazilian Journal of Biology</i> , 74 (4),		

Field Code Changed

647 **810-820**.

Barros Nogueira F. M., Silveira R. M. L., Girard P., Da Silva C. J., Abdo M. S. A. & 649 Wantzen K.M. 2011. Hydrochemistry of lakes, rivers and groundwater. In: Junk W. J., 650 Silva C. J., Cunha C. N., Wantzen K. M. (eds.) The Pantanal: Ecology, biodiversity 651 652 and sustainable management of a large neotropical seasonal wetland. Pensoft Publishers, Sofia, 167-198. 653 654 Barbiéro L., Queiróz-Neto J. P., Ciornei G., Sakamoto, A. Y., Cappelari B., Fernandes 655 E., & Valles V. 2002. Geochemistry of water and groundwater in the Nhecolândia, 656 Pantanal of Mato Grosso, Brazil: variability and associated process. Wetlands, 22(3), 657 528-540. 658 659 660 Barbiero L., Rezende Filho A., Furguim S. A. C., Furian A. Y., Sakamoto A. Y., Valles V., Graham R. C., Fort M., Ferreira R. P. D. & Queiroz Neto J. P. 2008. Soil 661 662 morphological control on saline and freshwater lake hydrogeochemistry in the Pantanal of Nhecolândia, Brazil. Geoderma, 148(1), 91-106 663 664 665 Batista D. S. N., Abreu U. G. P.; Feraz Filho P. B. & Rosa, A. N. 2012. Índices reprodutivos do rebanho Nelore da fazenda Nhumirim, Pantanal da Nhecolândia. 666 667 Acta Scientiarum Animal Science, 34(1), 71-76. 668 Becker B. F., da Silva-Caminha S. A. F., Guerreiro R. L., de Oliveira E. J., D'Apolito C. 669 & Assine, M. L. 2017. Late Holocene palynology of a saline lake in the Pantanal of 670 Nhecolândia, Brazil. Palynology, 1-9. 671 672

Bergier I. 2013. Effects of highland land-use over lowlands of the Brazilian Pantanal. *Science of the Total Environment*, **463**, 1060-1066.

675

676 677 678 679	Bergier I., Krusche A., & Guérin F. 2016. Alkaline lake dynamics in the Nhecolândia landscape. <i>In</i> : Bergier I. & Assine M. (eds) <i>Dynamics of the Pantanal Wetland in South America.</i> The Handbook of Environmental Chemistry, 37 , Springer, Cham, 145-161.
 680 681 682 683 684 	Bergier I., Assine M. L., McGlue M. M., Alho C. J., Silva A., Guerreiro R. L. & Carvalho, J. C. 2018. Amazon rainforest modulation of water security in the Pantanal wetland. <i>Science of The Total Environment</i> , 619 , 1116-1125.
685 686 687 688	Bertaux J., Sondag F., Santos R., Soubies F., Causse C., Plagnes V., Le Cornec F. & Seidel, A. 2002. Paleoclimatic record of speleothems in a tropical region: study of laminated sequences from a Holocene stalagmite in central-west Brazil. <i>Quaternary International</i> , 89 , 3-16.
689 690 691	Bespalez E. 2015. Arqueologia e história indígena no Pantanal. <i>Estudos Avançados</i> , 29 (83), p.45–86.
692 693 694 695	Blair T. C. & McPherson J. G. 1994. Alluvial fan processes and forms. <i>In:</i> Abrahams A. D. & Parsons A. J. (eds.) <i>Geomorphology of desert environments</i> . Chapman and Hall, London, 354-402.
696 697 698 699	Breeze P. S., Groucutt H. S., Drake N. A., White T. S., Jennings R. P. & Petraglia M. D. 2016. Palaeohydrological corridors for hominin dispersals in the Middle East~ 250–70,000 years ago. <i>Quaternary Science Reviews</i> , 144 , 155-185.
 700 701 702 703 704 	Buehler H. A., Weissmann G. S., Scuderi L. A. & Hartley, A. J. 2011. Spatial and temporal evolution of an avulsion on the Taquari River distributive fluvial system from satellite image analysis. <i>Journal of Sedimentary Research</i> , 81 (8), 630-640.

Bull W. B. & Schick A. P. 1979. Impact of climatic change on an arid watershed, 705 706 Nahal Yael, southern Israel. Quaternary Research, 11, 153-171. 707 708 Clapperton C. 1993. Quaternary Geology and Geomorphology of South America. Elsevier, Amsterdam. 709 710 Collischonn W., Tucci C. E. M. & Clarke, R. T. 2001. Further evidence of changes in 711 712 the hydrological regime of the River Paraguay: part of a wider phenomenon of 713 climate change? Journal of Hydrology, 245(1), 218-238. 714 715 Costa M. P. & Telmer K. H. 2006. Utilizing SAR imagery and aquatic vegetation to 716 map fresh and brackish lakes in the Brazilian Pantanal wetland. Remote sensing of 717 Environment, 105(3), 204-213. 718 719 Costa M. P. & Telmer K. H. 2007. Mapping and monitoring lakes in the Brazilian 720 Pantanal wetland using synthetic aperture radar imagery. Aquatic Conservation: 721 Marine and Freshwater Ecosystems, 17(3), 277-288. 722 723 Coutininho A. C. 2005. Dinâmica das queimadas no estado do Mato Grosso e suas 724 relações com as atividades antrópicas e a economia local. PhD thesis. Universidade 725 de São Paulo. 726 727 Evans T. L. & Costa M. 2013. Landcover classification of the Lower Nhecolândia subregion of the Brazilian Pantanal Wetlands using ALOS/PALSAR, RADARSAT-2 728 and ENVISAT/ASAR imagery. Remote Sensing of Environment, 128, 118-137. 729 730 Fantin-Cruz I., Pedrollo O., Castro N. M., Girard P., Zeilhofer P., & Hamilton S. K. 731 2011. Historical reconstruction of floodplain inundation in the Pantanal (Brazil) using 732

neural networks. *Journal of Hydrology*, **399**(3), 376-384.

Furian S., Martins E. R. C., Parizotto T. M., Rezende-Filho A. T., Victoria R. L. & 735 Barbiero L. 2013. Chemical diversity and spatial variability in myriad lakes in 736 737 Nhecolândia in the Pantanal wetlands of Brazil. Limnology and Oceanography, 58(6), 738 2249-2261. 739 Furquim S. A. C., Graham R. C., Barbiero L., de Queiroz N., Pereira J., & Vallès V. 740 741 2008. Mineralogy and genesis of smectites in an alkaline-saline environment of 742 Pantanal wetland, Brazil. Clays and Clay Minerals, 56(5), 579-595. 743 Furquim S. A. C., Graham R. C., Barbiéro L., Neto J. Q. & Vidal-Torrado P. 2010. Soil 744 mineral genesis and distribution in a saline lake landscape of the Pantanal Wetland, 745 Brazil. Geoderma, 154(3), 518-528. 746 747 748 Garcia E. A. C. 1984. O clima no Pantanal Mato-grossense. Circular Técnica 749 EMBRAPA, 14. 750 Guerreiro R. L., McGlue M. M. et al. 2017. Paleoecology explains Holocene chemical 751 changes in lakes of the Nhecolândia (Pantanal-Brazil). Hydrobiologia, 1-19. 752 753 Hamilton S. K., Sippel S. J. & Melack J. M. 2002. Comparison of inundation patterns 754 among major South American floodplains. Journal of Geophysical Research: 755 Atmospheres, 107(D20). 756 757 758 Harvey A. M. & Wells S. G. 1994. Late Pleistocene and Holocene changes in hillslope 759 sediment supply to alluvial fan systems: Zzyzx, California. In: Millington A. C. & Pye K. (eds) Environmental Change in Drylands: Biogeographical and Geomorphological 760

761 Perspectives. Wiley, Chichester, 67-84.

762

Horton B. K., & DeCelles P. G. 1997. The modern foreland basin system adjacent to 763 764 the Central Andes. Geology, 25(10), 895-898. 765 766 Huesemann, M. H. (2006). Can advances in science and technology prevent global warming? Mitigation and Adaptation Strategies for Global Change, 11(3), 539-577. 767 768 769 Klammer G. 1982. Die Paleowueste des Pantanal von Mato Grosso und die 770 pleistozaene Klimageschichte der brasilianischen Randtropen. Zeitschrift 771 Geomorphol NF, 26, 393-416 772 Kraus M. J. & Aslan, A. 1993. Eocene hydromorphic paleosols: Significance for 773 774 interpreting ancient floodplain processes. Journal of Sedimentary Petrology, 63, 775 453-463. 776 777 Krepper C. M., García N. O. & Jones P. D. 2006. Paraguay river basin response to 778 seasonal rainfall. International journal of climatology, 26(9), 1267-1278. 779 Lenters J. D. & Cook K. H. 1999. Summertime precipitation variability over South 780 America: role of the large-scale circulation. Monthly Weather Review, 127, 409-431 781 782 783 Harris M. B., Tomas W., Mourao G., Da Silva C. J., Guimarães E., Sonoda F. & Fachim E. 2005. Safeguarding the Pantanal wetlands: threats and conservation 784 785 initiatives. Conservation Biology, 19(3), 714-720. 786 Makaske B., Maathuis B. H. P., Padovani C. R., Stolker C., Mossleman E. & 787 Jongman, R. H. G. 2012. Upstream and downstream controls on recent avulsions on 788 789 the Taquari megafan, Pantanal, south-western Brazil. Earth Surface Processes and Landforms, 37(12), 1313-1326. 790

Mamede S. B. & Alho C. J. R. 2006. Response of wild mammals to seasonal 792 793 shrinking-and-expansion of habitats due to flooding regime of the Pantanal, Brazil. 794 Brazilian Journal of Biology, 66(4), 991-998. 795 Marengo J. A., Liebmann B., Kousky V. E., Filizola N. P., & Wainer I. C. 2001. Onset 796 797 and end of the rainy season in the Brazilian Amazon Basin. Journal of Climate, 14(5), 798 833-852. 799 Marengo J. A., Oliveira G. S. & Alves L. M. 2016a. Climate Change Scenarios in the 800 Pantanal. In: Bergier I. & Assine M. (eds) Dynamics of the Pantanal Wetland in South 801 America. The Handbook of Environmental Chemistry, 37, Springer, Cham, 227-238. 802 803 804 Marengo J. A., Alves L. M., & Torres R. R. 2016b. Regional climate change scenarios 805 in the Brazilian Pantanal watershed. Climate Research, 68(2-3), 201-213. 806 807 McGlue M. M., Silva A. et al. 2012. Lacustrine records of Holocene flood pulse 808 dynamics in the Upper Paraguay River watershed (Pantanal wetlands, Brazil). Quaternary Research, 78(2), 285-294. 809 810 811 McGlue M. M., Silva A., Assine M. L., Stevaux J. C. & do Nascimento Pupim F. 2015. 812 Paleolimnology in the pantanal: using lake sediments to track quaternary 813 environmental change in the world's largest tropical wetland. In: Bergier I. & Assine M. (eds) Dynamics of the Pantanal Wetland in South America. The Handbook of 814 Environmental Chemistry, 37, Springer, Cham, 51-81. 815 816

McGlue M. M., Guerreiro R. L., Bergier I., Silva A., Pupim F. N., Oberc V., & Assine M.
L. 2017. Holocene stratigraphic evolution of saline lakes in Nhecolândia, southern
Pantanal wetlands (Brazil). *Quaternary Research*, 1-19.

McGregor G. R. & Nieuwolt S. 1998. Tropical climatology: an introduction to the 821 822 climates of the low latitudes. John Wiley and Sons, London. 823 824 Meinshausen M., Meinshausen N., Hare W., Raper S. C., Frieler K., Knutti R., Frame D. J. & Allen M. R. (2009). Greenhouse-gas emission targets for limiting global 825 warming to 2 C. Nature, 458(7242), 1158-1162. 826 827 828 Nobre P., & Shukla J. 1996. Variations of sea surface temperature, wind stress, and 829 rainfall over the tropical Atlantic and South America. Journal of climate, 9(10), 2464-2479. 830 831 832 Oliveira E. C., Rossetti D. F. & Utida G. 2017. Paleoenvironmental Evolution of 833 Continental Carbonates in West-Central Brazil. Anais da Academia Brasileira de 834 Ciências, 89(1), 407-429. 835 836 Oliveira, J. E. 1994. A utilização da analogia etnográfica no estudo dos aterros da 837 região pantaneira de Corumbá. Revista de Arqueologia, 8, (2), 159-167. 838 Oliveira, J. E. 1996. Guató, argonautas do Pantanal. EDIPUCRS, Porto Alegre. 839 840 Padovani C. R., Pontara R. C. P. & Pereira J. G. 2001. Mudanças recentes de leito no 841 baixo curso do rio Taquari, no Pantanal Mato-grossense. Boletim de Geociências da 842 843 UFPR, 49, 33-38. 844 Paranhos Filho A., Moreira E., Oliveira A., Pagotto T. & Mioto, C. 2014. Análise da 845 variação da cobertura do solo no Pantanal de 2003 a 2010 através de sensoriamento 846 remoto. Engenharia Sanitária e Ambiental, 1(1), 69-76. 847

Paz A. R., Collischonn W., Bravo J. M., Bates P. D. & Baugh, C. 2014. The influence 849 of vertical water balance on modelling Pantanal (Brazil) spatio-temporal inundation 850 dynamics. Hydrological processes, 28(10), 3539-3553. 851 852 Pla-Pueyo S., Viseras C., Candy I., Soria J.M., García-García F. & Schreve D. 2015. 853 Climatic control on palaeohydrology and cyclical sediment distribution in the 854 855 Plio-Quaternary deposits of the Guadix Basin (Betic Cordillera, Spain). Quaternary International, 389, 56-69. 856 857 Por F. D. 1995. The Pantanal of Mato Grosso (Brazil): world's largest wetlands. 858 Volume 73. Kluwer Academic Publishers, Dordrecht. 859 860 Pott A. 1982. Pastagens das sub-regiões dos Paiaguás e da Nhecolândia do 861 Pantanal matogrossense. Circular Técnica EMBRAPA, 10. 862 863 864 Pott A. & Pott V. J. 1996. Flora of the Pantanal-Current listing of phanerogams (in Portuguese). In: Proceedings of the 2nd Symposium on Natural and Socio-Economic 865 Resources of the Pantanal, Corumbá, Brazil, 18–22 November 1996, 297–325. 866 867 868 Pott A. & Pott V. J. 2009. Vegetation of the Pantanal: Phytogeography and dynamics (in Portuguese). In: Proceedings of 2nd Symposium on Geo-Technologies of the 869 Pantanal, Corumbá, Brazil, 7–11, November 2009, 1065–1076. 870 871 872 Pott V. J., Pott A., Lima L. C. P., Moreira S.N. & Oliveira A.K.M. 2011a. Aquatic 873 macrophyte diversity of the Pantanal wetland and upper basin. Brazilian Journal of Biology, 71(Suppl.), 255-263. 874 875 876 Pott A. & Oliveira A. K. M. 2011b. Damasceno, G.A., Junior; Silva, J.S.V. 2011b.Plant diversity of the Pantanal wetland. Brazilian Journal of Biology, 71 (Suppl.), 265-273. 877

878

Ratter J. A., Pott A., Pott V. J., Cunha C. D. & Haridasan, M. 1988. Observations on
woody vegetation types in the Pantanal and at Corumbá, Brazil. *Notes from the Royal Botanic Garden*, 45,

882

Rezende-Filho A.T., Furian S., Victoria R.L., Mascré C., Valles V. & Barbiero, L. 2012.
Hydrochemical variability at the Upper Paraguay Basin and Pantanal wetland. *Hydrology and Earth Systems Science*, **16**(3),503–525.

886

Rodela L. G., Queiroz Neto J. P. & Santos, S. A. 2007. Classificação das pastagens
nativas do Pantanal da Nhecolândia, Mato Grosso do Sul, por meio de imagens de
satélite. *XIII Simpósio Brasileiro De Sensoriamento Remoto*, 4187-4194.

890

Rosa A. N., Abreu U. D. A., Silva L. O. C., Nobre P. R. C. & Gondo, A. P. R. C. V.
2007. Pecuária de corte no Pantanal brasileiro: realidade e perspectivas futuras de
melhoramento. *Embrapa Pantanal-Documentos*, 93.

894

Rossetto O. C. & Brasil Junior A. C. P. 2003. Cultura e desenvolvimento sustentável
no pantanal mato-grossense: entre a tradição e a modernidade. *Sociedade e Estado*, **18**(1-2), 155-175.

898

Sabino J. & Prado, P. I. 2006. Síntese do Conhecimento da Diversidade Biológica de
Vertebrados do Brasil. *In:* Levinsohn, T. (org.). *Avaliação do Estado do Conhecimento da Diversidade Brasileira*, 2, Ministério do Meio Ambiente, Brasília, 55-143.

902

903 Sakamoto A. Y., Bacani V. M., Gradella F. S., Ferreira C. C. & Decco H. F. 2012.
904 Desmatamento e alterações ambientais no Pantanal da Nhecolândia, MS, Brasil.
905 *Revista Geonorte - Edição Especial*, 3(4), 827-839.

Schmitz P. I., Rogge J. H., Beber M. V. & Rosa O. A. 2001. Arqueologia do Pantanal 907 908 do Mato Grosso do Sul - Projeto Corumbá. Tellus, 1(1), 11-26. 909 910 Schmitz P. I., Rogge J. H., Rosa A. O., Beber M. V. & Freitas E. A. V. 2009. Aterros 911 da tradição pantanal nas fazendas sagrado coração de Jesus e Bodoquena, Corumbá, 912 MS. Pesquisas, Antropologia, 67, 321-374. 913 914 Scremin-Dias E., Lorenz-Lemke A. P. & Oliveira A. K. M. 2011. The floristic 915 heterogeneity of the Pantanal and the occurrence of species with different adaptive 916 strategies to water stress. Brazilian Journal of Biology, 71 (Suppl.), 275-282. 917 918 Silva J. D. S. V., Abdon M. M., Pott A. & Mauro, R. A. 2009. Fragile ecosystem: the 919 Brazilian Pantanal wetland. Area studies-Brazil, 302. 920 921 Silva J., Abdon M. D. M., Silva S. M. A. & de Moraes, J. A. 2011. Evolution of 922 deforestation in the Brazilian Pantanal and surroundings in the timeframe 1976-2008. 923 Geografia (Rio Claro), 36(Especial), 35-55. 924 925 Slingerland R. L., & Smith N. D. 2004. River avulsions and their deposits. Annual Review of Earth and Planetary Sciences, 32, 257-285. 926 927 Soares A. P., Soares P. C. & Assine M. L. 2003. Areiais e lagoas do Pantanal, Brasil: 928 929 herança paleoclimática? Revista Brasileira de Geociências, 33, 211-224. 930 Thompson L. G., Mosley-Thompson E., & Henderson K. A. 2000. Ice-core 931 palaeoclimate records in tropical South America since the Last Glacial Maximum. 932 933 Journal of Quaternary Science, 15(4), 377-394.

935 Tricart J. 1982. El Pantanal: un ejemplo del impacto geomorfologico sobre el 936 ambiente. Investigaciones Geográficas, 29, 81. 937 938 Tubelis D. P. & Tomás W. M. 2002. Caracterização da avifauna da planicie do 939 Pantanal. Indicadores da magnitude da diversidade e abundância de vertebrados 940 silvestres do Pantanal num mosaico de hábitats sazonais. EMBRAPA, Corumbá. 941 942 Úbeda B., Di Giacomo A. S., Neiff J. J., Loiselle S. A., Poi A. S. G., Galvez J. A., & 943 Cozar A. 2013. Potential effects of climate change on the water level, flora and macro-fauna of a large Neotropical Wetland. PloS one, 8(7), e67787. 944 945 946 Ussami N., Shiraiwa S., & Dominguez J. M. L. 1999. Basement reactivation in a 947 sub-Andean foreland flexural bulge: The Pantanal wetland, SW Brazil. Tectonics, 948 18(1), 25-39. 949 950 Valeriano M. M., Salvi L. L. & Aragão J. R. L. 2012. Relações entre a distribuição da 951 precipitação e o relevo da bacia do alto Paraguai. IV Simposio de Geotecnologias no 952 Pantanal, 289-298 953 954 Vera C., Higgins W., et al. 2006. Toward a unified view of the American monsoon 955 systems. Journal of climate, 19(20), 4977-5000. 956 957 Vialou D., Benabdelhadi M., Feathers J., Fontugne M. & Vialou A. V. 2017. Peopling South America's centre: the late Pleistocene site of Santa Elina. Antiquity, 91, 958 865-884. 959 960 Vitousek P. M. 1994. Beyond 961 global warming: ecology and global 962 change. Ecology, 75(7), 1861-1876. 963

964 Weyler G. 1962. *Relatório final dos poços perfurados no Pantanal* 965 *Matogrossense-Projeto Pantanal*. Petrobras, Ponta Grossa.

966

Whitney B. S., Mayle F. E. *et al.* 2011. A 45kyr palaeoclimate record from the lowland
interior of tropical South America. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **307**(1), 177-192.

970

Zani H., Assine M. L. & McGlue M. M. 2012. Remote sensing analysis of depositional
landforms in alluvial settings: Method development and application to the Taquari
megafan, Pantanal (Brazil). *Geomorphology*, **161**, 82-92.

974

975 Figure captions

Fig.1. Location of the Pantanal basin. Circles highlighting Nhecolândia and Pantanal
do Miranda regions, over a digital elevation model from Shuttle Radar Topographic
Mission (SRTM).

Fig.2. Comparison between rainfall and Paraguay river levels in the same station(Porto Manga, Corumbá), without correlation (ANA 2017).

Fig.3. Landsat 7 image of the Taquari megafan (acquired April, 2000) displaying
historical and active lobes of the alluvial fan (numbered 1 to 7 and outlined by white
dashed lines, modified from Makaske et al. 2012). The currently active depositional
lobe is outlined in solid white lines. The location of the incised meander belt is
highlighted, as is the Nhecolândia region.

986 Fig.4. Aerial photographs showing a freshwater (baía or lagoa, depending on the authors describing them) and a saline lake (salina) in the Nhecolandia region (taken 987 from Google Earth, 2017) and sketches showing their cross-section during the wet 988 989 and dry seasons, following the descriptions by Barbiero et al. 2008 and Furiam et al. 990 2013. These authors propose the existence of chemically cemented soils below the saline lakes which prevent freshwater from reaching the salinas during the dry 991 992 season, when evaporation is higher. On the other hand, during the flood events in the wet season, the salinity is reduced through the fresh groundwater reaching the lake. 993



Figure 2

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Comparison - Porto Manga - River level vs Rainfall



Active depositional fan lobe

6

Incised meander belt

Coxim

100

3

5

O

2

Nhecolândia

